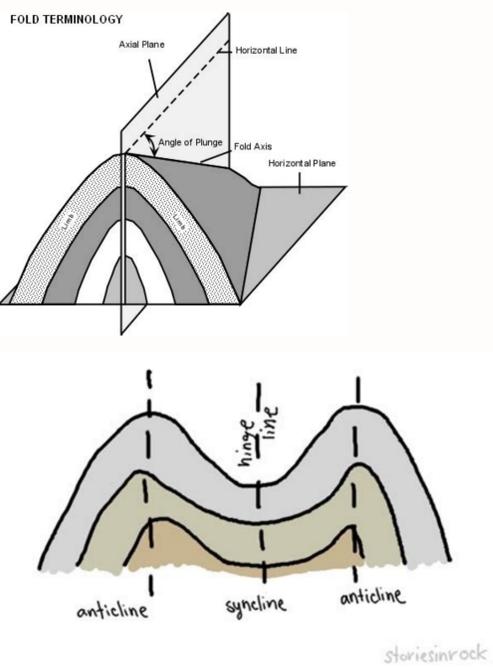
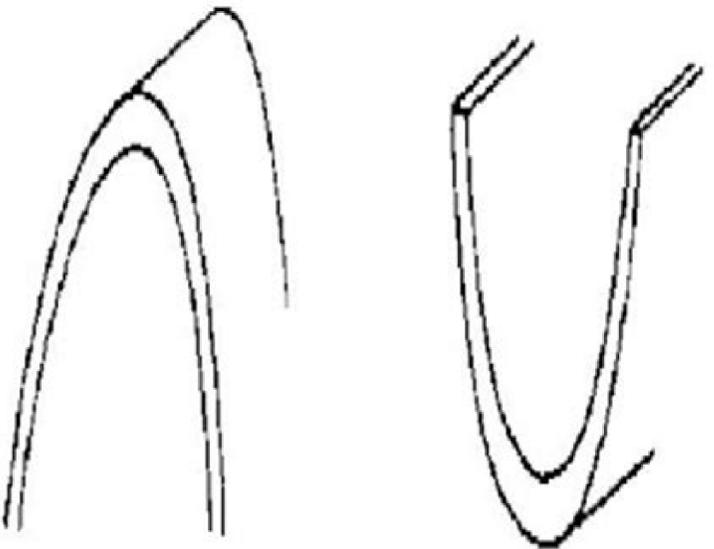
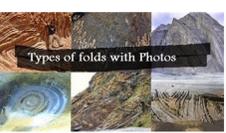


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antiform | anticline | Antiform is a holonym of anticline. In context[geology]lang=en terms the difference between antiform and anticline is that antiform is (geology) a topographic feature which is composed of sedimentary layers in a convex formation, but may not actually form a real anticline (ie , the oldest rocks may not be exposed in the middle) while anticline is (geology) a fold with strata sloping downwards on each side. As nouns the difference between antiform and anticline is that antiform is (geology) a topographic feature which is composed of sedimentary layers in a convex formation, but may not actually form a real anticline (ie , the oldest rocks may not be exposed in the middle) while anticline is (geology) a fold with strata sloping downwards on each side. (en noun) (geology) A topographic feature which is composed of sedimentary layers in a convex formation, but may not actually form a real anticline (ie , the oldest rocks may not be exposed in the middle). (art) A form that defies the usual conventions of artistic forms. * 1983 , Edith Kurzweil, William Phillips, Writers and Politics: A Partisan Review Reader (page 236) Dada forms and antiforms were used as a means of reacting against Abstract Expressionism and abstract art generally (en noun) (geology) A fold with strata sloping downwards on each side. This Diagram depicts some of the differences between Asymmetrical, Symmetrical, and OVERTURNED folds. Credit: W. H. Freeman and Company A wave-like geologic structure that forms when rocks deform by bending instead of breaking under compressional stress. Anticlines are arch-shaped folds in which rock layers are upwardly convex. The oldest rock layers form the core of the fold, and outward from the core progressively younger rocks occur. A syncline is the opposite type of fold, having downwardly convex layers with young rocks in the core. Folds typically occur in anticline-syncline pairs. The hinge is the point of maximum curvature in a fold. The limbs occur on either side of the fold hinge. The imaginary surface bisecting the limbs of the fold is called the axial surface. The axial surface is called the axial plane in cases where the fold is symmetrical and the lines containing the points of maximum curvature of the folded layers, or hinge lines, are coplanar. Concentric folding preserves the thickness of each bed as measured perpendicular to original bedding. Similar folds have the same wave shape, but bed thickness changes throughout each layer, with thicker hinges and thinner limbs. Folds form under varied conditions of stress, hydrostatic pressure, pore pressure, and temperature gradient, as evidenced by their presence in soft sediments, the full spectrum of metamorphic rocks, and even as primary flow structures in some igneous rocks. A set of folds distributed on a regional scale constitutes a fold belt, a common feature of orogenic zones. Folds are commonly formed by shortening of existing layers, but may also be formed as a result of displacement on a non-planar fault (fault bend fold), at the tip of a propagating fault (fault propagation fold), by differential compaction or due to the effects of a high-level igneous intrusion e.g. above a laccolith. Fold Classification Folds are classified on the basis of several geometric factors: Tightness of folding The tightness of folds can be described as open (limbs dip gently), tight (limbs dip steeply) or isoclinal (limbs are parallel). Orientation of axial plane The orientation of the axial plane relative to the horizontal together with the orientation of fold limbs allow subdivision into upright (axial plane vertical, limbs symmetric), overturned (axial plane moderately inclined, one limb overturned), or recumbent (axial plane near horizontal, one limb inverted). Thickness of folded beds Thickly-bedded, brittle units tend to form concentric folds with the bed thickness preserved normal to bedding surfaces. Thinly-bedded, clay-rich units have a tendency to develop a foliation parallel to the axial plane and form similar folds with the vertical distance between top and bottom of the unit preserved through the deformation. Anticline: linear, strata normally dip away from axial center, oldest strata in center. Syncline: linear, strata normally dip toward axial center, youngest strata in center. Antiform: linear, strata dip away from axial center, age unknown, or inverted. Synform: linear, strata dip toward axial center, age unknown, or inverted. Dome: nonlinear, strata dip away from center in all directions, oldest strata in center. Basin: nonlinear, strata dip toward center in all directions, youngest strata in center. Monocline: linear, strata dip in one direction between horizontal layers on each side. Chevron: angular fold with straight limbs and small hinges Recumbent: linear, fold axial plane oriented at low angle resulting in overturned strata in one limb of the fold. Slump: typically monoclinical, result of differential compaction or dissolution during sedimentation and lithification. Ptygmatic: Folds are chaotic, random and disconnected. Typical of sedimentary slump folding, migmatites and decollement detachment zones. Parasitic: short wavelength folds formed within a larger wavelength fold structure - normally associated with differences in bed thickness Disharmonic: Folds in adjacent layers with different wavelengths and shapes Recommended For You Oldest ice core ever drilled outside the polar regions Barstow syncline in Rainbow Basin, Mojave Desert, California Photo Copyright © Garry Hayes In structural geology, an anticline is a type of fold that is an arch-like shape and has its oldest beds at its core. A typical anticline is convex up in which the hinge or crest is the location where the curvature is greatest, and the limbs are the sides of the fold that dip away from the hinge. Anticlines can be recognized and differentiated from antiforms by a sequence of rock layers that become progressively older toward the center of the fold. Therefore, if age relationships between various rock strata are unknown, the term antiform should be used. The progressing age of the rock strata towards the core and uplifted center, are the trademark indications for evidence of Anticlines on a geologic map. These formations occur because Anticlinal ridges typically develop above thrust faults during crustal deformations. The uplifted core of the fold causes compression of strata that preferentially erodes to a deeper stratigraphic level relative to the topographically lower flanks. Motion along the fault including both shortening and extension of tectonic plates, usually also deforms strata near the fault. This can result in an asymmetrical or overturned fold. Terminology of anticlines Anticline at Calico Ghost Town Location: San Bernardino County, California, United States. Photo Copyright © Garry Hayes An Antiform can be used to describe any fold that is convex up. It is the relative ages of the rock strata that separate anticlines from antiforms. The hinge of an anticline refers to the location where the curvature is greatest, also called the crest. The hinge is also the highest point on a stratum along the top of the fold. The culmination also refers to the highest point along any geologic structure. The limbs are the sides of the fold that display less curvature. The inflection point is the area on the limbs where the curvature changes direction. The axial surface is an imaginary plane connecting the hinge of each layer of rock stratum through the cross sectional anticline. If the axial surface is vertical and the angles on each side of the fold are equivalent, then the anticline is symmetrical. If the axial plane is tilted or offset then the anticline is asymmetrical. An anticline that is cylindrical has a well-defined axial surface, whereas non-cylindrical anticlines are too complex to have a single axial plane. Formation processes Anticlines are usually developed above thrust faults, so any small compression and motion within the inner crust can have large effects on the upper rock stratum. Stresses developed during mountain building or during other tectonic processes can similarly warp or bend bedding and foliation (or other planar features). The more the underlying fault is tectonically uplifted, the more the strata will be deformed and must adapt to new shapes. The shape formed will also be very dependent on the properties and cohesion of the different types of rock within each layer. Recommended For You Fire Opal : What Is Fire Opal? How Is Fire Opal Formed? During the formation of flexural-slip folds, the different rock layers form parallel-slip folds to accommodate for buckling. A good way to visualize how the multiple layers are manipulated, is to bend a deck of

cards and to imagine each card as a layer of rock stratum. The amount of slip on each side of the anticline increases from the hinge to the inflection point. Passive-flow folds form when the rock is so soft that it behaves like weak plastic and slowly flows. In this process different parts of the rock body move at different rates causing shear stress to gradually shift from layer to layer.iii There is no mechanical contrast between layers in this type of fold. Passive-flow folds are extremely dependent on the rock stratum's makeup and can typically occur in areas with high temperatures Syncline In structural geology, a syncline is a fold with younger layers closer to the center of the structure. A synclinalorium (plural synclinaloriums or synclinaloria) is a large syncline with superimposed smaller folds. Synclines are typically a downward fold, termed a synformal syncline (i.e. a trough); but synclines that point upwards, or perched, can be found when strata have been overturned and folded (an antiformal syncline). Dome A dome is a feature in structural geology consisting of symmetrical anticlines that intersect each other at their respective apices. Intact, domes are distinct, rounded, spherical-to-ellipsoidal-shaped protrusions on the Earth's surface. However, a transect parallel to Earth's surface of a dome features concentric rings of strata. Consequently, if the top of a dome has been eroded flat, the resulting structure in plan view appears as a bullseye, with the youngest rock layers at the outside, and each ring growing progressively older moving inwards. These strata would have been horizontal at the time of deposition, then later deformed by the uplift associated with dome formation. Basin A structural basin is a large-scale structural formation of rock strata formed by tectonic warping of previously flat lying strata. Structural basins are geological depressions, and are the inverse of domes. Some elongated structural basins are also known as synclines. Structural basins may also be sedimentary basins, which are aggregations of sediment that filled up a depression or accumulated in an area; however, many structural basins were formed by tectonic events long after the sedimentary layers were deposited. Basins appear on a geologic map as roughly circular or elliptical, with concentric layers. Because the strata dip toward the center, the exposed strata in a basin are progressively younger from outside-in, with the youngest rocks in the center. Basins are often large in areal extent, often hundreds of kilometers across. Structural basins are often important sources of coal, petroleum, and groundwater. Monocline A monocline (or, rarely, a monoform) is a step-like fold in rock strata consisting of a zone of steeper dip within an otherwise horizontal or gently-dipping sequence. Chevron Chevron folds are a structural feature characterized by repeated well behaved folded beds with straight limbs and sharp hinges. Well developed, these folds develop repeated set of v-shaped beds. They develop in response to regional or local compressive stress. Inter-limb angles are generally 60 degrees or less. Chevron folding preferentially occurs when the bedding regularly alternates between contrasting competences. Turbidites, characterized by alternating high-competence sandstones and low-competence shales, provide the typical geological setting for chevron folds to occur. Perpetuation of the fold structure is not geometrically limited. Given a proper stratigraphy, chevrons can persist almost indefinitely. Chevron folds with flat-lying axial planes, Millook Haven, North Cornwall, UK Credit: Smalljim Recumbent An asymmetrical fold is one in which the axial plane is inclined. An overturned fold, or overfold, has the axial plane inclined to such an extent that the strata on one limb are overturned. A recumbent fold has an essentially horizontal axial plane. When the two limbs of a fold are essentially parallel to each other and thus approximately parallel to the axial plane Recumbent fold at Godrevy in Cornwall in England. The rocks are of Devonian age and they were folded during the Variscan orogeny. Credit: mwcarruthers Slump Typically monoclinal, result of differential compaction or dissolution during sedimentation and lithification. Slump Fold An almost isoclinal fold (coin, seaweed and shells for scale) formed as wet layers of mud settled and solidified in Triassic times. Credit: Anne Burgess Folds are chaotic, random and disconnected. Typical of sedimentary slump folding, migmatites and decollement detachment zones. Ptygmatic folding, Broken Hill Credit: Monash University Short wavelength folds formed within a larger wavelength fold structure - normally associated with differences in bed thickness Disharmonic Folds in adjacent layers with different wavelengths and shapes Flexural slip Flexural slip allows folding by creating layer-parallel slip Folding mechanisms between the layers of the folded strata, which, altogether, result in deformation. A good analogy is bending a phone book, where volume preservation is accommodated by slip between the pages of the book. The fold formed by the compression of competent rock beds is called "flexure fold". Buckling Typically, folding is thought to occur by simple buckling of a planar surface and its confining volume. The volume change is accommodated by layer parallel shortening the volume, which grows in thickness. Folding under this mechanism is typically of the similar fold style, as thinned limbs are shortened horizontally and thickened hinges do so vertically. Mass displacement If the folding deformation cannot be accommodated by flexural slip or volume-change shortening (buckling), the rocks are generally removed from the path of the stress. This is achieved by pressure dissolution, a form of metamorphic process, in which rocks shorten by dissolving constituents in areas of high strain and redepositing them in areas of lower strain. Folds created in this way include examples in migmatites, and areas with a strong axial planar cleavage. Photos

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